

6.1 Mag. Quake, Hakui, Japan - 11/26/2024

M 6.1 - 35 km NW of Hakui, Japan
 2024-11-26 13:47:03 UTC | 37.072°N 136.467°E | 10.0 km depth



Administrative Region

Data not available.

Nearby Places

Place	Distance	Population
Hakui, Ishikawa, Japan	35 km (21.8 mi) SE	Population: 24,529
Anamizu, Ishikawa, Japan	42.4 km (26.3 mi) ENE	Population: 17,842
Nanto, Ishikawa, Japan	44.4 km (27.6 mi) E	Population: 43,320
Tsubata, Ishikawa, Japan	50.8 km (31.6 mi) SSE	Population: 13,154
Kanazawa, Ishikawa, Japan	54.3 km (33.8 mi) SSE	Population: 462,225

Distance and direction from epicenter to nearby place. Nearby place information was automatically generated from [OpenStreetMap](#). This list includes populated places, not necessarily cities. If this information is inaccurate, please consider [update the Geonames database](#).

Tectonic Summary

Seismotectonics of Japan and Vicinity

The North America plate, Pacific plate, Philippine Sea plate, and Eurasia plate all influence the tectonic setting of Japan, Taiwan, and the surrounding area. Some authors divide the edges of these plates into several microplates that together take up the overall relative motions between the larger tectonic blocks, including the Okhotsk microplate in northern Japan, the Okinawa microplate in southern Japan, the Yangtze microplate in the area of the East China Sea, and the Amur microplate in the area of the Sea of Japan.

The seafloor expression of the boundary between the Pacific and North America plates lies 300 km off the east coasts of Hokkaido and Honshu at the Kuril-Kamchatka and Japan trenches. The subduction of the Pacific plate beneath the North America plate, at rates of 83-100 mm/yr, generates abundant seismicity, predominantly as a result of interplate slip along the interface between the plates. The 1952 M 8.4 Torofu, 1963 M 8.5 Kuri, 2003 M 8.3 Tokachi-Oki, and the 2011 M 9.0 Tohoku earthquakes all eventually reach megathrust seismicity. The 1933 M 8.4 Sanriku-Oki earthquake and the 1994 M 8.3 Shikotan earthquake are examples of intraplate seismicity, caused by deformation within the lithosphere of the subducting Pacific plate (Sanriku-Oki) and of the overriding North America plate (Shikotan), respectively.

At the southern terminus of the Japan Trench the intersection of the Pacific, North America, and Philippine Sea plates forms the Boso Triple Junction, the only example of a trench-trench intersection in the world. South of the triple junction the Pacific plate subducts beneath the Philippine Sea plate at the Izu-Ogasawara trench, at rates of 45-56 mm/yr. This margin is noteworthy because of the steep dip of the subducting Pacific plate (70° or greater below depths of 50 km depth), and because of its heterogeneous seismicity; few earthquakes above M 7 occur at shallow depths, yet many occur below 400 km. The lack of large shallow megathrust earthquakes may be a result of weak coupling at the plate interface, or simply a reflection of an incomplete earthquake catalog with respect to the length of typical seismic cycles.

The northernmost section of the Philippine Sea plate shares a 350 km boundary with the North America plate that runs approximately east-west from the Boso Triple Junction towards the Izu Peninsula. This short boundary is dominated by the subduction of the Philippine Sea plate beneath Japan along the Sagami Trough, but also includes small sections of transform motion.

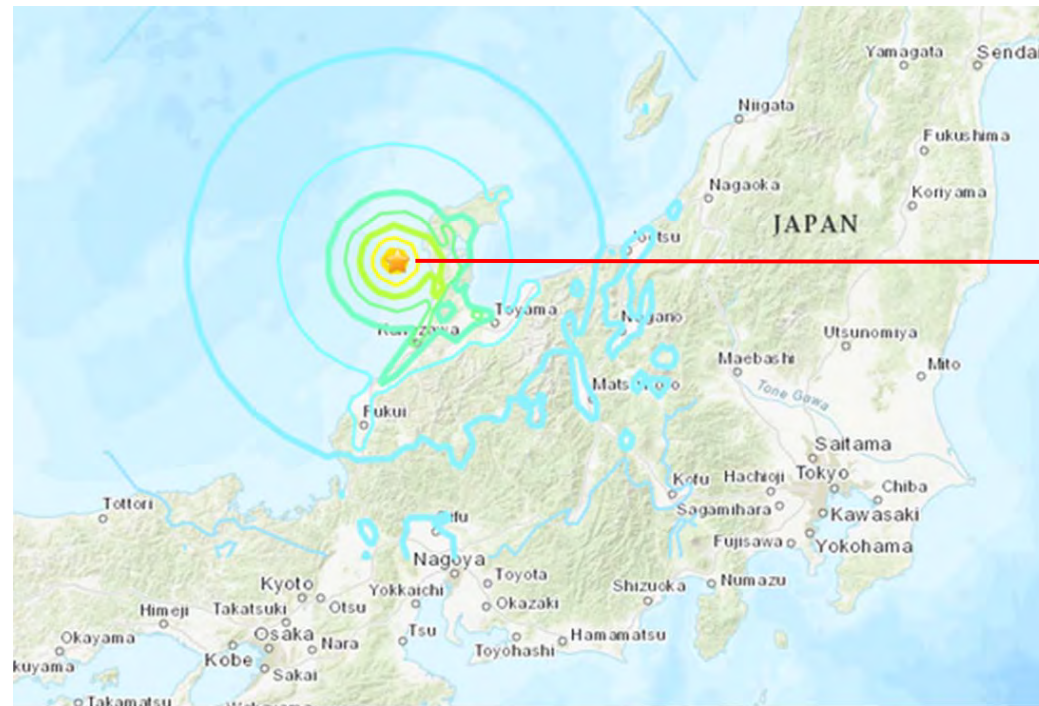
The subduction of the Philippine Sea plate under the Eurasia plate begins at the Suruga Trough, immediately southwest of the Izu peninsula. In the northern Tokai, Tonankai and Nankai sections of this subduction zone, historical data indicate M 8+ earthquake recurrence intervals of 100-150 years. The Tonankai and Nankai sections last ruptured in M 8.1 earthquakes in 1944 and 1946, respectively, while the Tokai section last broke in 1854. In the 1980's studies began to forecast the imminence of a large earthquake in the Tokai region, and warned of its potential impact on the cities of Tokyo and Yokohama (the two largest cities in Japan) to date, the expected event has not occurred.

The boundary between the Philippine Sea and Eurasia plates continues south and southwestwards from the Suruga Trough, extending 2000 km along the Nankai and Ryukyu trenches before reaching the island of Taiwan. Along the Ryukyu Trench, the Philippine Sea plate exhibits trench-normal subduction at rates increasing from 46 mm/yr in the northeast to 65 mm/yr in the southwest. Convergence and the associated back-arc deformation west of the oceanic trench creates the Ryukyu islands and the Okinawa Trough. The largest historic event observed along this subduction zone was the M 8.1 Kikai Island earthquake in 1911.

In the vicinity of Taiwan the structure of the Philippine Sea-Eurasia plate boundary and the associated pattern of seismicity becomes more complex. 400 km east of Taiwan a clockwise rotation in the trend of the margin (from NE-SW to E-W), paired with an increase in subduction obliquity creates a section of the plate boundary that exhibits dextral transform and oblique thrusting motions. South of Taiwan the polarity of subduction flips; the Eurasia plate subducts beneath the Philippine Sea plate. Debate surrounds contrasting models of the plate boundary position between the zones of oppositely verging subduction, and the boundary's relation to patterns of seismicity. Many studies propose that crustal thickening causes the majority of regional seismicity, while others attribute seismicity to deformation associated with subduction. Another resolution proposes a tear in the Philippine Sea plate and a complex assortment of subduction, transform, and collisional motion. All the models concede that seismicity around the island of Taiwan is anomalously shallow, with few earthquakes deeper than 70m.

While there are no instances of an earthquake M 8 in the modern record, Taiwan and its surrounding region have experienced eight M 7-7.5 events between 1900 and 2014. The dominance of shallow M 8 earthquakes suggests fairly weak plate boundary coupling, with most earthquakes caused by internal plate deformation. The 1935 M 7.1 Hsinchu-Tsiching earthquake and the 1999 M 7.4 Chi-Chi Earthquake both eventually reach shallow continental crust thrust faulting that dominates regional seismicity across the island. A major tectonic feature of the island is the Longitudinal Valley Fault, which ruptures frequently as small, shallow earthquakes. In 1951, the Longitudinal Valley Fault hosted twelve M 6.5 events known as the Hualien-Taipei earthquake sequence.

Large earthquakes in the vicinity of Japan and Taiwan have been both destructive and deadly. The regions high population density makes shallow earthquakes especially dangerous. Since 1900 there have been 13 earthquakes (9 in Japan, 4 in Taiwan) that have each caused over 1000 fatalities, leading to a total of nearly 200,000 earthquake related deaths. In January 1995 an earthquake that ruptured a southern branch of the Japan Median Tectonic Line near the city of Kobe (population 1.5 million) killed over 5000 people. The 1923 Kanto earthquake shook both Yokohama (population 500,000, at that time) and Tokyo (population 2.1 million), killing 142,000 people. The earthquake also started fires that burned down 90% of the buildings in Yokohama and 40% of the buildings in Tokyo. Most recently, the M 9.0 Tohoku earthquake, which ruptured a 400 km stretch of the subduction zone plate boundary east of Honshu, and the tsunami it generated caused over 20,000 fatalities.

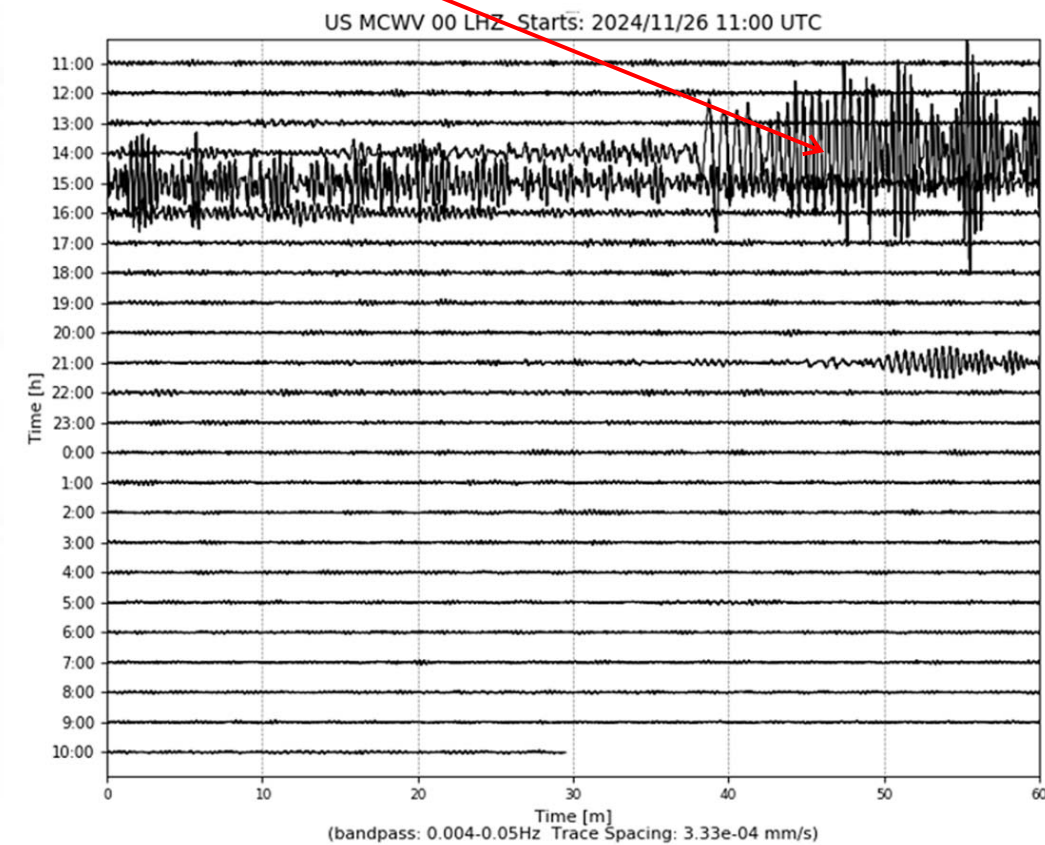
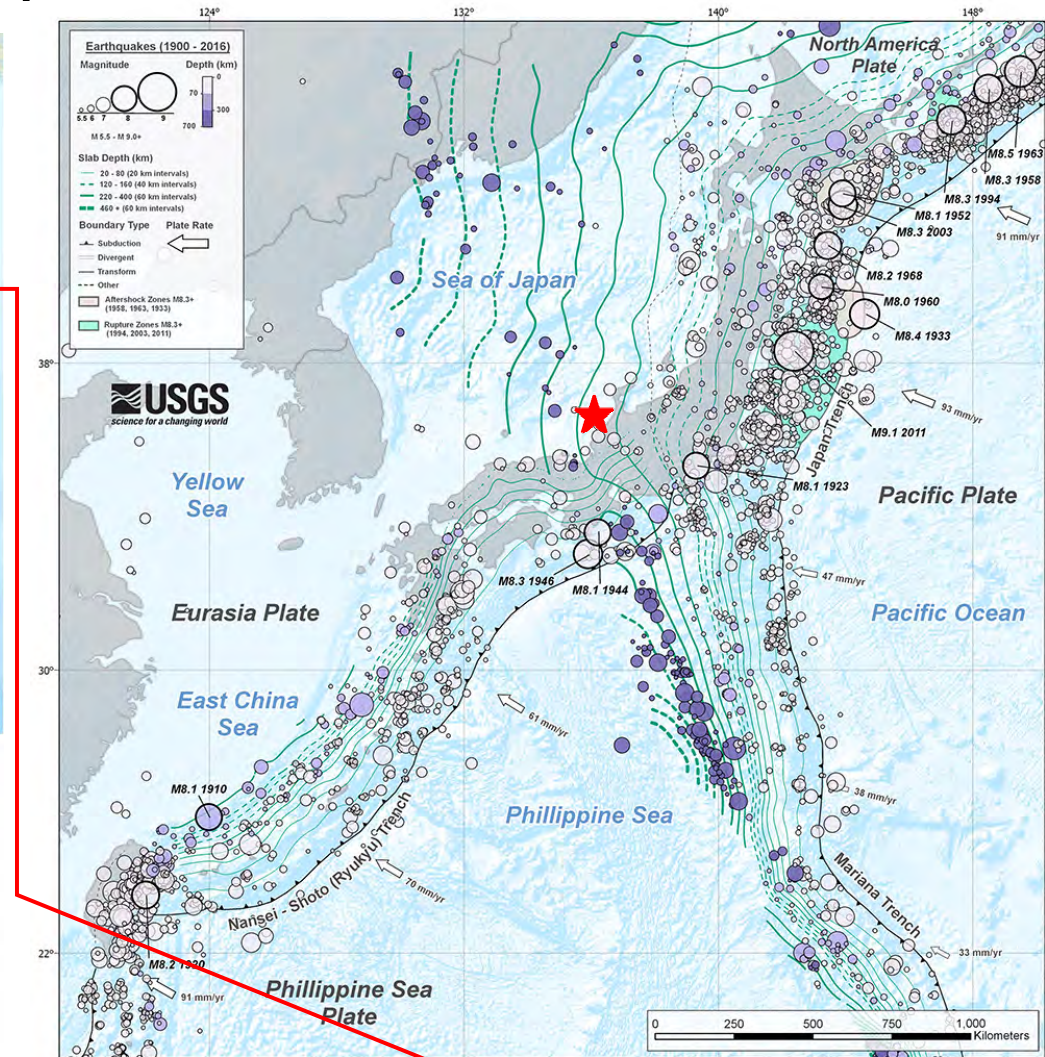


USGS Community Internet Intensity Map
 NEAR THE WEST COAST OF HONSHU, JAPAN
 2024-11-26 13:47:03 UTC 37.0718N 136.4671E M6.1 Depth: 10 km ID:us7000nu90



SHAKING	Not felt	Weak	Light	Moderate	Strong	Very strong	Severe	Violent	Extreme
DAMAGE	none	none	none	Very light	Light	Moderate	Moderate/Heavy	Heavy	Very Heavy
INTENSITY	I	II-III	IV	V	VI	VII	VIII	IX	X+

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Large (6.1 Mag.), shallow (10 Km) offshore quake. No tsunami.